

# METHODS FOR THE CATION EXCHANGE CAPACITY OF SALINE-SODIC SOILS<sup>1</sup>

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## ABSTRACT

Cation exchange capacities of four soil profiles representing large ranges in salinity and Na saturation were measured by: 1) the sum of cations ( $CEC_s$ ) extracted by  $NH_4OAc$  at pH 7 minus the cations in saturation extracts, 2) the method of Bower et al (1952) ( $CEC_b$ ) in which the soil is saturated using  $NaOAc$ , the excess salt is removed by washing with ethanol and the Na is removed with  $NH_4OAc$ , and 3) the method of Polemio and Rhoades (1977) ( $CEC_{PR}$ ) in which the soil is saturated with Na using a solution of  $NaCl-NaOAc$  in 60% ethanol, the Na is replaced with  $Mg(NO_3)_2$  and the Cl concentration in the  $Mg(NO_3)_2$  solution is used to correct occluded salt. Values for CEC were essentially the same for all methods in the non-saline and non-sodic Yolo soil and in the surface horizons of the other three soils which were relatively free of salts. However, in the moderately to lightly saline subsoils of the Pescadero, Solano and Pescadero Variant soils, the  $CEC_s$  greatly exceeded the  $CEC_b$  and  $CEC_{PR}$ . Even though the  $CEC_b$  and  $CEC_{PR}$  generally gave the same values, the  $CEC_b$  values were more highly correlated with hygroscopic and saturation water contents. Using these correlations as criteria for a satisfactory method, the  $CEC_b$  appears to be the most reliable of the three methods that were compared.

**Keywords:** Dissolution of minerals, Hygroscopic water content, Sum of cations.

## COMPENDIO

Se determinó la Capacidad de Intercambio Catiónico de cuatro perfiles de suelos representativos de un amplio rango de salinidad y saturación de sodio medida por: 1) La suma de cationes ( $CIC_s$ ) extraídos con  $NH_4OAc$  a pH 7.0 corregidos por cationes en solución, 2) El método de Bower ( $CIC_b$ , 1952) en el cual el suelo se satura con  $NaOAc$ , el exceso de sal se remueve mediante lavado con etanol y el sodio desplazado con  $NH_4OAc$  y 3) El método de Polemio y Rhoades ( $CIC_{PR}$ , 1977) en el cual el suelo se satura con Na usando una solución de  $NaCl-NaOAc$  en etanol al 60%, el Na se desplaza luego con  $Mg(NO_3)_2$  y la concentración de Cl en esta última solución se usa para corregir la sal ocuida. Los valores de la  $CIC$  resultaron esencialmente iguales en el suelo Yolo no salino y no sódico y en los horizontes superficiales de los otros tres suelos. (relativamente libres de sales). Sin embargo, en la moderada a fuertemente salinos horizontes subsuperficiales de los suelos Pescadero, Solano y Pescadero Variant, la  $CIC_s$  excede notablemente la  $CIC_b$  y la  $CIC_{PR}$ . Aún cuando la  $CIC_b$  y la  $CIC_{PR}$  dan generalmente los mismos valores, los valores de la  $CIC_b$  correlacionaron más estrechamente con los contenidos de agua higroscópica y con el agua de saturación. Usando dichas correlaciones como criterio de selección, la  $CIC_b$  parece ser el método más confiable de los tres comparados.

**Palabras claves:** Dilución de minerales, Contenido higroscópico de agua, Suma de cationes

## INTRODUCTION

The cation-exchange capacity (CEC), defined as the quantity of readily exchangeable cations neutralizing negative charge in the soil, is an important property of

soils. The CEC is dependent on the conditions under which it is measured and it varies with the method used (Chapman, 1965). Ideally the method for determining CEC should be selected to fit the specific soil, but in

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many laboratories a standard procedure for all soil samples is essential for rapid routine analyses (Rhoades, 1982).

Problems of dissolution of minerals and fixation of cations during CEC determinations have been discussed by Chapman and Kelley (1930), Bower (1950, 1955), Bower et al (1952), Babcock (1960), Rhoades and Krueger (1968), Polemio and Rhoades (1977) Sayegh et al (1978), and Rhoades (1982). Methods for alkaline, saline, calcareous, gypsiferous and sodic soils have been proposed by Bower et al (1952), Shatemirov (1960), Peech et al (1962), Vergina and Dokuchayev (1964), Matushevsky (1968), Carpena et al (1972), Papanicolaou (1976), and Polemio and Rhoades (1977).

The objective of the research reported herein was to determine the CEC values of four soil profiles representing extreme ranges in salinity, exchangeable cations and other soils properties. The sum of cations extracted by  $\text{NH}_4\text{OAc}$  was used to determine its limitations and the magnitude of overestimation of the CEC in saline soils. The Bower et al. (1952) method and the Polemio and Rhoades (1977) methods were compared because they have both been recommended for alkaline saline soils.

## METHODS AND MATERIALS

### Soil Samples

Samples of the horizons of four soil profiles were collected to represent a wide range in salinity and sodicity. The samples were air dried and ground to pass a 2-mm sieve. The profiles were the Pescadero, Pescadero Variant, Yolo and Solano.

The Pescadero soil was classified as a fine montmorillonitic thermic Aquic Natrixeralf. The pH was alkaline and increased with an increase in depth. The organic matter content ranged from 5.8% in the upper layer to 0.79% in the lower layers. The profile was slightly saline at the surface with an electrical conductivity of  $3.3 \text{ dS m}^{-1}$  which increased to  $4.3 \text{ dS m}^{-1}$  at the lower horizons. The profile was a silty clay soil to a depth of 73 cm. The dominant minerals in the clay fraction were montmorillonite and kaolinite. Because of high saturation with Na, the soil is dedicated to pasture and has not been irrigated.

The Solano soil was acid to a depth of 26 cm but the rest of the profile was alkaline with an increase in pH with increase in depth to a maximum pH of 8.5 for the lowest horizon. The organic matter decreased with increased depth from 2.44% in the upper layer to 0.26% in the lower layers. The profile was nonsaline to a depth of 42 cm and sodic to a depth of 93 cm. The clay content

ranged from 11.2% at the surface layer to a maximum of 38.0% for the B22t (42-64 cm) horizon dominated by montmorillonite, kaolinite and vermiculite. The profile was classified as a loamy mixed thermic Typic Natrargid. This soil is under pasture and has not been irrigated.

The Yolo soil was classified as a fine silty mixed non-acid thermic Typic Xerorthent. It is a silt loam to a depth of 85 cm with a fine sandy layer between 85 and 108 cm. The dominant clay minerals were montmorillonite, kaolinite and vermiculite. It was an alkaline, nonsaline and nonsodic soil with organic matter content ranging from 2.4% to 0.26%. This soil has been irrigated and cultivated for a long time and has evidently been well leached.

The Pescadero Variant soil (fine montmorillonitic thermic Aquic Natrargid) was situated in a depression of the field which is flooded during the rainy season. The pH ranged from neutral to highly alkaline. It was highly saline with electrical conductivities ranging from  $7.7$  to  $28.3 \text{ dS m}^{-1}$  and also it had a high Na saturation. This soil is so impermeable that it is unproductive. The average clay content was 34.7%, dominated by montmorillonite and kaolinite.

### General Methods

Water lost by drying air-dried soils at  $105^\circ\text{C}$  was used as a measure of hygroscopic water ( $P_w$ ). The water content of a saturated paste (U.S. Salinity Laboratory Staff, 1954) was used as a measure of water content at saturation (SP). Values for pH were determined on the saturated paste and the electrical conductivity of the saturation extract,  $\text{dS m}^{-1}$ , was used as a measure of soil salinity. The organic matter (OM) was determined by a Walkley-Black procedure (Chapman and Pratt, 1961). Texture analyses were determined using the method of Day (1965).

### Cation-Exchange Capacity

**Sum of Cations ( $\text{CEC}_s$ ).** A 5-g sample was placed in a centrifuge tube and 33-ml aliquot of  $\text{M NH}_4\text{OAc}$  solution adjusted to pH 7, was added. The soil suspension was shaken for 5 minutes, centrifuged at  $\text{RCF} = 1000$  until the supernatant liquid was clear and then the supernatant was decanted into a 100-ml volumetric flask. This extraction was repeated a total of three times. The volume of the extract was diluted to 100 ml and the Ca, Mg, Na and K were determined by atomic absorption spectrophotometry (AAS) and flame spectrophotometry (FS). The sum of cations extracted by  $\text{NH}_4\text{OAc}$  minus the sum of soluble cations in the saturation extract was used as an estimate of the CEC.

**Bower Method ( $\text{CEC}_b$ ).** A 4-g sample of soil was placed in a centrifuge tube and a 33-ml aliquot of  $\text{M}$

NaOAc solution, adjusted to pH 8.2, was added. The soil suspension was shaken for 5 minutes and then centrifuged at RCF = 1000 until the supernatant liquid was clear. The supernatant was then decanted and discarded. The sample was extracted in this manner a total of four times. The sample was then washed four times with 95% ethanol to remove the excess NaOAc solution. The adsorbed Na was replaced by three extractions with 33-ml aliquots of  $M$   $NH_4OAc$  solution adjusted to pH 7, and the extracted Na was determined by FS (Bower et al., 1952).

**Polemio and Rhoades Method ( $CEC_{PR}$ ).** The method of Polemio and Rhoades (1977) was developed for arid land soils, particularly for those containing carbonates, gypsum and zeolites.

A 4.0-g sample of soil was saturated with Na by four successive equilibrations with 33-ml aliquots of a 60% ethanol solution containing 0.4  $M$  NaOAc-0.1  $M$  NaCl adjusted to pH 8.2. During each equilibration the soil suspension was shaken for 5 minutes, centrifuged until the supernatant was clear after which the supernatant liquid was discarded. The Na-saturated samples plus the occluded saturating solution were extracted three

times with 33-ml aliquots of a 0.5  $M$  pH 7 solution of  $Mg(NO_3)_2$ . Total Na and Cl were determined in the extract. The exchangeable Na was calculated as the total Na minus the occluded soluble Na which was 5 times the occluded Cl.

## RESULTS AND DISCUSSION

The relationships between salinity, expressed in  $dS\ m^{-1}$ , and depth for four soil profiles are presented in *Figura 1*. The Yolo soil was free of salt through the profile. The Solano soil was nonsaline in the surface but increased to moderate levels of salinity in the subsoil. The Pescadero profile was moderately saline through the first meter and decreased in salinity at deeper depths. The Pescadero Variant soil was saline at all depths and highly saline at depths lower than about 35 cm. For the four profiles the salinity ranged from less than 1 to 28  $dS\ m^{-1}$ , providing excellent soil materials for testing the reliability of methods for determining CEC.

Data for  $NH_4OAc$ -extracted cations, the sum of cations ( $CEC_s$ ), the  $CEC_b$  and  $CEC_{PR}$ , all plotted as a function of profile depth, are pre-sented in *Figs. 2 to 5*, respectively, for the Pescadero, Solano, Yolo and Pescadero Variant soils. The Pescadero soil was definitely sodic to a depth of about 60 cm with much lower amounts of extracted Na below that depth. Extracted Mg exceeded extracted Ca throughout most of the profile. The average  $CEC_b$  and  $CEC_{PR}$  values were essentially the same throughout the profile, but the SC gave a large overestimation of the CEC except in the surface layers. Based on the profiles for extracted cations the overestimation of the CEC by SC resulted from extraction of Ca and Mg from the dissolution of soil minerals in  $NH_4OAc$ . The extracted Ca exceeded  $CEC_b$  and  $CEC_{PR}$  for the deepest depth and the extracted Mg was equal to the average of the  $CEC_b$  and  $CEC_{PR}$  at one depth. As the difference between  $CEC_s$  and  $CEC_b$  increased the extracted Na decreased, indicating indirectly that dissolution of Na minerals might have contributed very little to the excess of  $CEC_s$  over  $CEC_b$ .

For the Solano soil the three methods for determining CEC gave essentially the same values in the three surface horizons. The first two horizons of this soil were slightly acid and the third horizon was only slightly alkaline. All of these surface horizons were relatively free of salt. Below the 42-cm depth the  $CEC_s$  exceeded the  $CEC_b$  with a dominant effect of Ca and Mg dissolution in the deepest horizons.

The Yolo Soil contained very little extractable Na. The extracted Mg exceeded the extracted Ca by a small amount throughout the profile. The  $CEC_s$  agreed well

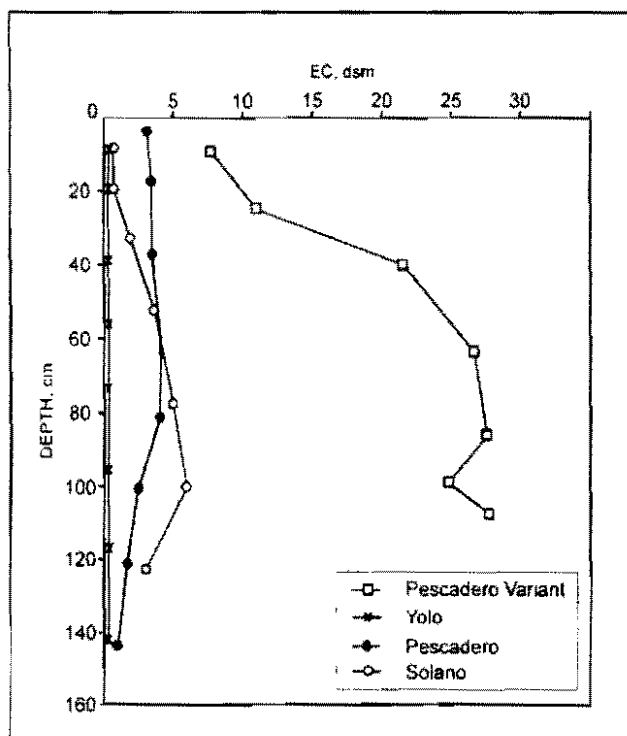


Figure 1. Salinity profiles for four soil profiles of northern California

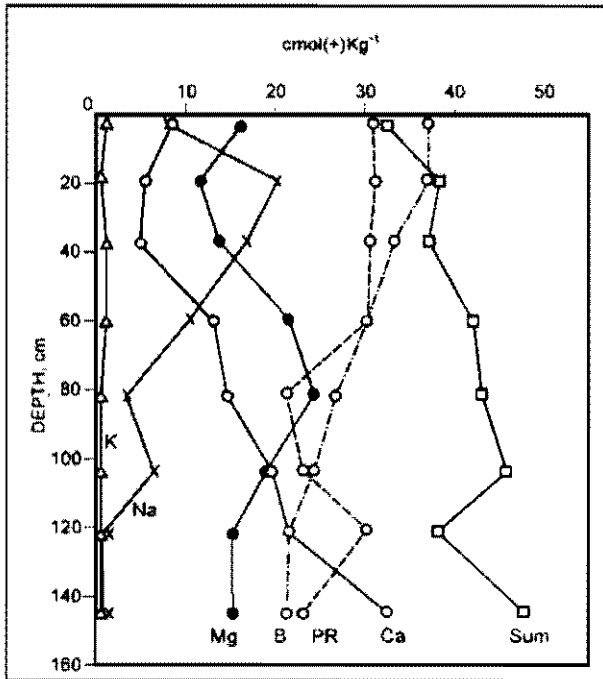


Figure 2. Relationship among the CEC values determined by three methods and the exchangeable cations in the Pescadero Soil

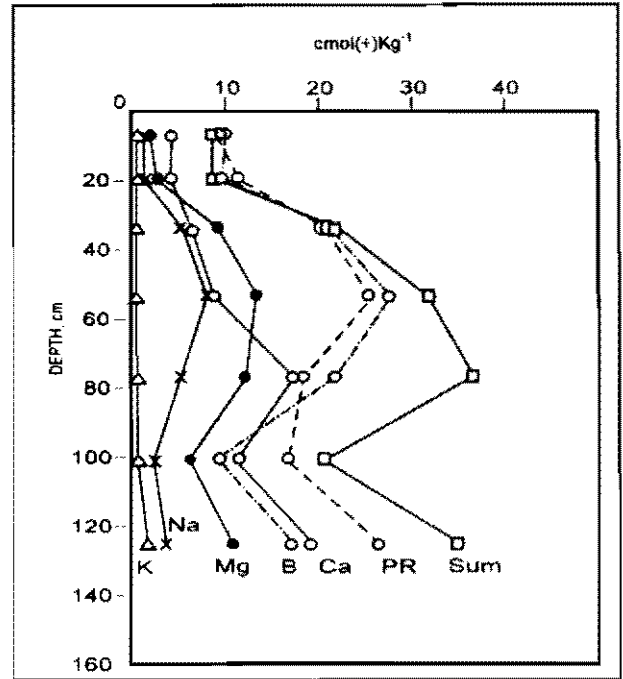


Figure 3. Relationship among the CEC values determined by three methods and the exchangeable cations in the Solano Soil

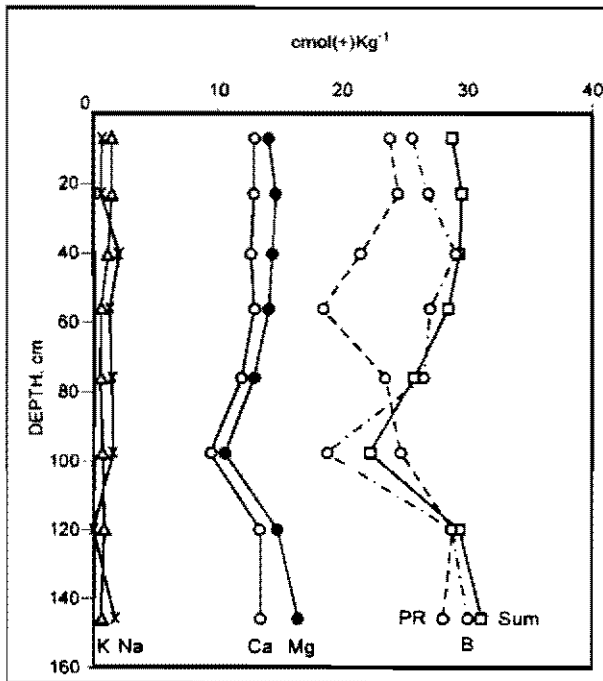


Figure 4. Relationship among the CEC values determined by three methods and the exchangeable cations in the Yolo Soil

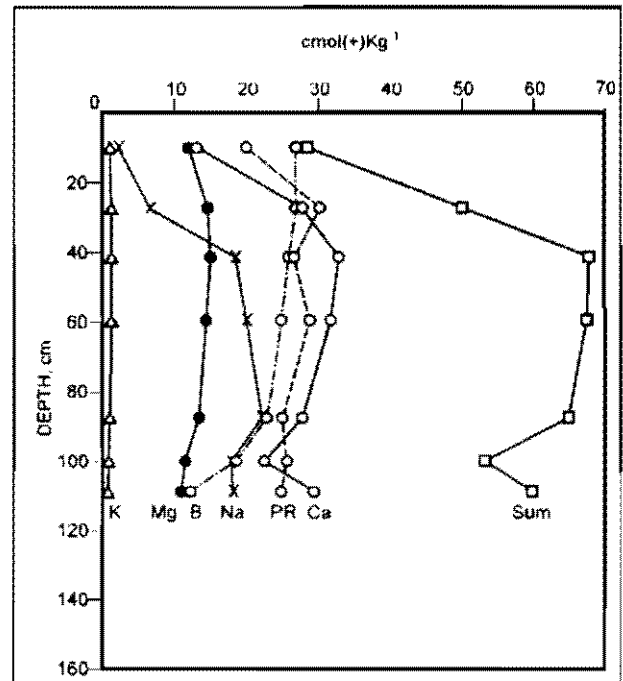


Figure 5. Relationship among the CEC values determined by three methods and the exchangeable cations in the Pescadero Variant Soil

with the  $CEC_B$  but the  $CEC_{PR}$  was less than the other  $CEC$  values for several horizons. The data indicate very little if any dissolution of minerals in  $NH_4OAc$  in this soil indicating that the extracted cations are exchangeable cations. This soil is nonsaline nonsodic and because of the consistency and agreements among the  $CEC$  values, particularly the  $CEC_S$ , it indicates that the methods were appropriately used and that the data for the other three saline-sodic profiles are real and not a result of errors.

For the Pescadero Variant soil the  $CEC_S$  was nearly equal to the  $CEC_B$  in the surface horizon, but it greatly exceeded the  $CEC_B$  for all other depths. The distribution of cations suggests that the dissolution of Ca and Na were likely the cause of high  $CEC_S$  values in this profile. Extracted Ca exceeded  $CEC_B$  and  $CEC_{PR}$  in all except the surface horizon. Extracted Na was correlated with  $CEC_S$  whereas extracted Mg was nearly constant throughout the profile. This saline-sodic soil provides an extreme case of dissolution of minerals in  $NH_4OAc$  solution and is an excellent example of the overestimation of  $CEC$  by the SC extracted in  $NH_4OAc$ .

Correlation coefficients for the relationships of  $CEC$  values with  $P_w$  and SP for individual soil profiles are presented in Table 1. The  $CEC_B$  was highly correlated with  $P_w$  for each soil, whereas the  $CEC_{PR}$  was correlated with  $P_w$  for only one soil. The  $CEC_{PR}$  was not correlated with any. The correlation and regression coefficients for the relationships of  $CEC_B$  with organic matter (OM) and clay contents are presented in Table 2. There were no significant relationships with OM but the  $CEC_B$  was related to clay contents for each soil.

**Table 1. Correlation coefficients for the linear relationships of  $CEC$  values with soil water contents of air dry soils ( $P_w$ ) and saturated soils (SP)**

SOIL	$CEC_B$	$CEC_{PR}$
<b>Pw</b>		
Pescadero	0.986 ***	0.656 ns
Solano	0.990 ***	0.821 *
Yolo	0.991 ***	0.101 ns
Pescadero variant	0.889 *	0.400 ns
All soils	0.802 ***	0.729 ***
<b>Sp</b>		
Pescadero	0.750 *	0.562 ns
Solano	0.933 ***	0.712 ns
Yolo	0.061 ns	0.015 ns
Pescadero variant	0.536 ns	0.630 ns
All soils	0.497 **	0.400 *

ns, \*, \*\*, and \*\*\*, respectively, indicate nonsignificant at the 95% level and significant at the 95%, 90% and 99.9% levels of probability

Pescadero = fine montmorillonitic thermic Aquic Natrizeralf.

Solano = loamy mixed thermic Typic Natrargid.

Yolo = fine silty mixed nonacid thermic Typic Xerorthent.

Pescadero Variant = fine montmorillonitic thermic Aquic Natrargid.

**Table 2. Correlation and regression coefficients for  $CEC_B$  values with soil organic matter and clay contents**

PROPERTY	SOIL	SLOPE	INTERCEPT	r
Organic matter	Pescadero	0.319	22.90	0.782 ns
	Solano	- 0.384	20.20	0.432 ns
	Yolo	0.166	2.08	0.294 ns
	Pescadero Variant	0.432	19.20	0.582 ns
Clay	Pescadero	0.050	11.10	0.930 **
	Solano	0.068	0.37	0.989 **
	Yolo	0.067	11.90	0.843 *
	Pescadero Variant	0.060	- 0.95	0.882 **

ns Non-significant value  
 \* Significant at the 5% level of significance  
 \*\* Significant at the 1% level of significance

When multiple correlation was used to evaluate the relation of clay content, organic matter and the water content of the saturated paste, using data for all profiles, the best correlation and regression was:

$$Y = 5.19 + 0.0342X_1 + 0.0233X_2^{1/4}$$

with  $R^2 = 0.654$  and  $R = 0.0805$

where,  $Y = \text{CEC}_B$  in  $\text{cmol}(+) \text{kg}^{-1}$   
 $X_1 = \text{clay content in g Kg}^{-1}$   
 $X_2 = \text{OM in g Kg}^{-1}$

The water content of the saturated paste was not significant in multiple correlation and regression largely because it was highly correlated with  $X_1$ , the clay content, with  $r = 0.84$ .

The multiple correlation coefficient  $R^2 = 0.654$  means that 65% of the variation in the  $\text{CEC}_B$  was associated with clay and organics matter content. The other 35%

was probably related to the  $\text{CEC}_S$  of the clay fractions of the four soils. The  $\text{CEC}$  per  $\text{Kg}$  of clay average 75, 70, 124 and 57  $\text{cmol}(+) \text{Kg}^{-1}$ , respectively, for the Pescadero, Solano, Yolo and Pescadero Variant soils.

According to Banin and Amiel (1969) two soil properties, the saturation percentage (SP) and air-dry hygroscopy water content (Pw), are proportional to  $\text{CEC}$ . Polemio and Rhoades (1977) used these characteristics to test indirectly the accuracy of their  $\text{CEC}$  method since there exists no absolute reference method. Their coefficients of linear correlation showed better correspondence between  $\text{CEC}$  and SP, and between  $\text{CEC}$  and Pw when determined by the newly developed  $\text{CEC}_{PR}$  method. The same procedure was used to evaluate the results of the present research. The better correspondence was found between the  $\text{CEC}_B$  and Pw which gave positive  $r$  values significant at the 1% level. The  $\text{CEC}_{PR}$  gave a significant  $r$  value of 0.82 for only the Solano soil. Also the  $\text{CEC}_{PR}$  was not correlated with SP for any soil.

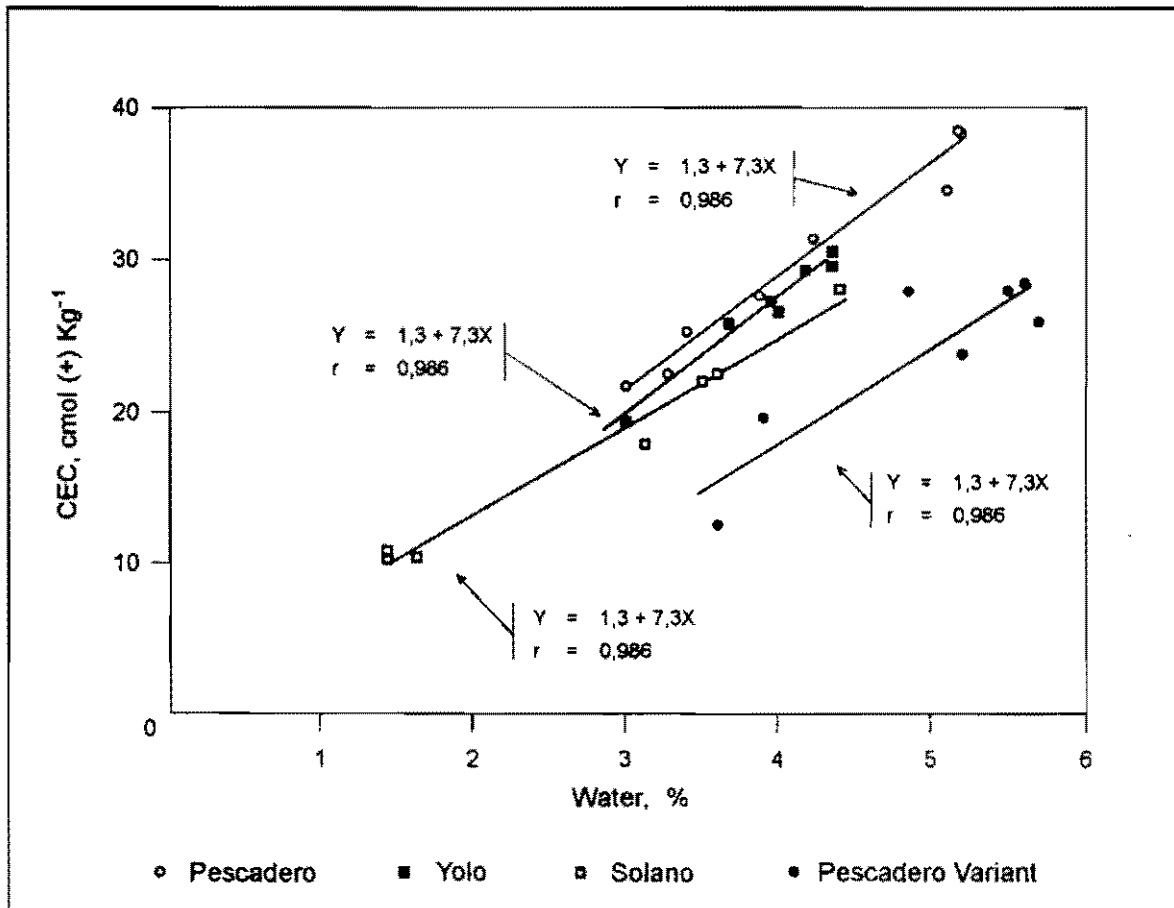


Figure 6. Relationship between  $\text{CEC}_B$  and the hygroscopic water content for each of four soils

The relationship between  $CEC_b$  and hygroscopic water content,  $P_w$ , for each soil presented in Fig. 6. The regression equations for the Pescadero, Solano and Yoio are essentially the same and the  $CEC_b$  values at 4% water have a narrow range of 24 to 28  $cmol(+) kg^{-1}$ . The data for the Pescadero Variant soil, however, show a lower correlation coefficient and a lower CEC per unit of hygroscopic water. This soil also showed the lowest CEC per unit of clay. The data for this soil also explain the lower correlation coefficients for all soils as compared to the coefficients for each soil (Table 1).

The lack of correspondence between the  $CEC_{PR}$  and the two soil properties for all the soils except for  $P_w$  in Solano soil, support the conclusion that this method is not adequate for the CEC determination in the soils studied. Similarly, given the high correlation between CEC and  $P_w$  for all the soils, and between CEC and SP for Pescadero and Solano soils when the  $CEC_b$  values are used, given credence to the conclusion that the method of Bower et al. (1952) is the most reliable for saline-sodic soils.

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